Modelling CO2 storage in large-scale aquifer systems

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Mathematics and Cybernetics, SINTEF Digital Joint with: O. Andersen, H. M. Nilsen, R. Allen, and O, Møyner

The 20th European Conference on Mathematics for Industry 18-22 June 2018, Budapest, Hungary Most IEA/IPCC scenarios require *widespread use of CCS* to reach the 2° target, latest scenario requires storing 3800 Mt/yr

Even if renewable energy grows 5% annually

- continuing current situation will give 3.2° warming
- large-scale CCS will be largely responsible to pick up the slack (through negative emissions)

"Three years to safeguard our climate" (Nature, June 2017)
the *entire* carbon budget might be spent in just 4 years, or 15 years on average



(2005) Storage, IPCC and Capture oon Dioxide Special Report Adapted from:

Methods for storing CO2 in deep underground geological formations



Thinking big

Sleipner: first offshore CO $_2$ storage project. Injected 1 Mt/yr into the Utsira formation since 1996



Thinking big

Sleipner: first offshore CO $_2$ storage project. Injected $1 \rm Mt/yr$ into the Utsira formation since 1996

From Sleipner to central storage hub in the North Sea.

Need to store several gigatonnes per year to significantly affect European point emissions

This is more than 1000 times the magnitude of ongoing operations



Primary physical processes of CO_2 storage

- Injection driven by viscous pressure drop
- CO₂ forms a supercritical phase that is lighter and weakly soluble in water



- Buoyant plume will migrate upward in the formation, displacing brine
- Upward movement limited by the caprock





To evaluate feasibility of possible injections, one must assess:

- capacity: how much?
- injectivity: how fast?
- safety: will it leak?

Challenging task because site characteristics vary a lot and data are scarse



Illustration: Eiken et al, Lessons Learned from 14 years of CCS Operations: Sleipner, In Salah and Snøhvit, Energy Procedia (2011)

Simulation models

Somewhat simplified, consist of three parts:

- a geological model volumetric grid with cell/face properties describing the porous rock formation
- a flow model describes how fluids flow in a porous medium (conservation laws + appropriate closure relations)
- 3 a well model describes flow in and out of the reservoir, in the wellbore, flow control devices, surface facilities





Geologic model: representative elementary volume



Flow model: governing equations for fluid flow

Conservation of mass

$$\frac{\partial}{\partial t} \int_{V} m \, dx + \oint_{\partial V} \vec{F} \cdot \vec{n} \, ds = \int_{V} r \, dx$$

m = mass, $\vec{F} = flow$ rate, r = fluid sources



Flow model: governing equations for fluid flow

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Oarcy's law – empirical law for processes on unresolved scale

$$\vec{u} = -\mathbf{K}(\nabla p - \rho g \nabla z)/\mu$$



Similar to Fourier's law (heat) [1822], Ohm's law (electric current) [1827], Fick's law (concentration) [1855], except that we now have *two driving forces*

Flow model: single-phase flow

$$rac{\partial(\phi
ho)}{\partial t} +
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ight) = q, \qquad ec{u} = -rac{\mathbf{K}}{\mu} ig(
abla p -
ho g
abla zig)$$

Special cases:

incompressible flow:

 $-\nabla \cdot \left(\mathbf{K}\nabla p\right) = q$

weakly compressible flow:

$$\frac{\partial p}{\partial t} = \frac{1}{\mu\phi c} \nabla \cdot \left(\mathbf{K} \nabla p \right)$$



Flow model: multiphase flow



Multiphase extension of Darcy's law (Muskat, 1936):

$$\vec{u}_{\alpha} = -\frac{\mathbf{K}_{\alpha}(S_{\alpha})}{\mu_{\alpha}} \big(\nabla p_{\alpha} - \rho_{\alpha} g \nabla z \big),$$

Effective permeability experienced by one phase is reduced by the presence of other phases, $\mathbf{K}_{\alpha} = \mathbf{K} k_{r\alpha} (S_{\alpha_1}, \ldots, S_{\alpha_m}), \quad 0 \leq k_{r\alpha} \leq 1$

Mass-balance equations for each phase (Muskat, 1945):

$$\frac{\partial(\phi\rho_{\alpha}S_{\alpha})}{\partial t} + \nabla\cdot\left(\rho_{\alpha}\vec{u}_{\alpha}\right) = q_{\alpha}$$

Elliptic pressure equation

$$\nabla \cdot \vec{u} = q, \qquad \vec{u} + \lambda \mathbf{K} \left[\nabla p - (f_w \rho_w + f_n \rho_n) \vec{g} - f_w \nabla P_c \right] = 0$$

Hyperbolic transport equation (saturation equation)

$$\phi \frac{\partial S_w}{\partial t} + \nabla \left[f_w \vec{u} + f_w \lambda_n (\rho_w - \rho_n) \mathbf{K} \vec{g} + f_w \lambda_n P'_c \nabla S_w \right] = q_w.$$

Notation used above:	
$\lambda_{\alpha} = k_{k\alpha}/\mu_{\alpha}$	
$f_{\alpha} = \lambda_{\alpha} / \sum_{\alpha} \lambda_{\alpha}$	

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Four different physical effects

- ► advection —
- convection
- segregation
- capillarity

Notation used above: $\lambda_{lpha} = k_{klpha}/\mu_{lpha}$ $f_{lpha} = \lambda_{lpha}/\sum_{lpha}\lambda_{lpha}$

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Elliptic pressure equation

$$\begin{array}{ll} \nabla \cdot \vec{u} = q, & \vec{u} + \lambda \mathbf{K} \left[\ \nabla p \ - \ (f_w \rho_w + f_n \rho_n) \vec{g} \ - \ f_w \nabla P_c \end{array} \right] = 0 \\ \mbox{Hyperbolic transport equation (saturation equation)} \\ & \phi \frac{\partial S_w}{\partial t} + \nabla \left[\ f_w \vec{u} \ + \ f_w \lambda_n (\rho_w - \rho_n) \mathbf{K} \vec{g} \ + \ f_w \lambda_n P'_c \nabla S_w \end{array} \right] = q_w. \\ \mbox{Four different physical effects} \\ & \ advection \\ & \ convection \\ & \ segregation \\ & \ capillarity \end{array}$$

A true multiscale problem



Estimating storage capacity

Site-specific estimates

Dynamic estimates: direct numerical simulations of specific site

Can incorporate technical, economical and regulatory constraints



From: "CO₂ Storage Atlas, Norwegian North Sea" (NPD, 2011)

Computationally challenging to incorporate necessary geological uncertainty in 3D simulations

Regional estimates

Statistical/analytical approach; no simulation



From: "Carbon Utilization and storage Atlas 2012" (USDOE, NETL)'

Computationally intractable at necessary 3D resolution

Simplest possible analysis: structural trapping

Model reduction: eliminate time and only consider infinitesimal, buoyant trickle of CO_2

Simplest possible analysis: structural trapping



Model reduction: eliminate time and only consider infinitesimal, buoyant trickle of CO_2

Analogy with hydrology



- traps \rightarrow lakes
- spill paths \rightarrow rivers
- spill points \rightarrow lake outlets
- catchment areas \rightarrow catchment areas
- watersheds \rightarrow watersheds





Observations:

- very large aspect ratios
- large difference in fluid densities, which means buoyancy effects are strong
- gravity segregation is almost instantaneous compared with lateral flow
- CO₂ plume is very thin and hence high vertical resolution is required

Two natural simplifications:

- depth-integrate the equations
- assume vertical equilibrium



Incompressible flow:

$$b \frac{\partial s}{\partial t} + \nabla \cdot \left[\lambda \mathbf{K} (\nabla p - \rho \vec{g}) \right] = 0$$

Depth integration:

$$\int_{top}^{bottom} \left[\phi \frac{\partial s}{\partial t} + \nabla \cdot \left[\lambda \boldsymbol{K} \big(\nabla p - \rho \vec{g} \big) \right] \right] = 0$$

Assuming sharp interface:

$$\phi \frac{\partial h}{\partial t} + \nabla_{\parallel} \cdot \left[\lambda \boldsymbol{K} \big(\nabla_{\parallel} P - \rho \vec{g} \big) \right] = 0$$



Example: Johansen formation

 CO_2 saturation 500 yrs after start of injection (110 year injection period)

Numerical efficiency for VE:

- vertical equilibrium \longrightarrow only one cell in the vertical direction \longrightarrow much fewer grid cells
- weaker dynamic coupling \longrightarrow fewer time steps

Time-step restrictions for explicit scheme

Time	Advection	Convection	Segregation
injection	1 yr	201 yr	8 yr
migration		164 yr	8 yr
injection	0.1 yr	10 yr	0.04 yr
migration		12 yr	0.04 yr
VE mod	lel 3	D model	

Example: Johansen formation

Analytical expression for vertical fluid distribution \longrightarrow 'infinite' vertical resolution

In particular, more accurate than underresolved 3D simulation



General VE formulation

Formulated as industry-standard black-oil type model[†]:

- capillary fringe
- compressibility
- dissolution and convective mixing
- geological heterogeneity
- hysteretic effects
- subscale trapping effects
- geomechanical, thermal,
 Robust and fully implicit
 discretization



 \dagger Using relative permeability $k_{r\alpha}(S,p)$ upscaled (homogenized) to match behavior on subscale grid

Year 1



< 1%4%< 1%







Year 110

















Calibrating the Sleipner model



Calibrating the Sleipner model





 $\rm CO_2$ density change: 760 \rightarrow 478 $\rm kg/m^3$ lnjection rate change: 0.92 \times benchmark rates

Grid elevation change: +/- 5 meters Permeability change: $1.85 \rightarrow 12$ darcy Porosity change: $0.36 \rightarrow 0.37$

More complex geology \longrightarrow hybrid 3D–VE simulation

- Fully automated coarsening for models with many (near) impermeable horizons
- Regions automatically detected and discretized
- Near-well regions: 3D grid with coupling wells and surface facilities
- Transition between VE-zones, diffuse leakage, fine-scale are all included



Coupling types:



VE to VE



Fine-scale to VE

Example: Sleipner-type scenario (injection)



Example: Sleipner-type scenario (migration)



Optimization of injection points and rates

$$O(q) = \int_0^T \sum_i q_i \, dt - C \cdot L(q) - C_p \sum_j \left[\max(0, p_j - p_j^{\max}) \right]^2$$

Three contributions to objective function

- Injected amount of CO_2
- Amount of CO_2 leaked at end of simulation, penalized by factor C
- Cells with pressure p_j exceeding safety limit p_j^{max} , penalized by factor C_p

Remarks:

- Each evaluation of O requires a full simulation acellerate
- Our framework enables ∇O to be computed by an adjoint approach[†]
- For brevity, no constraints are specified here

† J.D. Jansen, Adjoint-based optimization of multi-phase flow through porous media - a review, Computers & Fluids 46 (2011)

Accelerated simulation: early exit by forecast algorithm





Optimization with multiple constraints

How much can we safely inject into this formation to maximize storage while maintaining the integrity of the caprock?



Well placement, initial rates, and trapping structure Overburden pressure (MPa)



Strategies to enchance storage

Relocating wells downslope from trap increases storage from 380 to 850 Mt



Adding two water producers increases storage from 380 to 574 Mt



- Discussed modelling of long-term CO₂ storage in large-scale aquifer systems
- Multiscale problem, computationally intractable with standard 3D simulation
- Presented various model reductions:
 - Spill-point/trapping analysis static capacity estimates
 - Vertical equilibrium: reduce dimension and improve dynamic couplings
- Discussed various optimization methods: initial guess from trapping analysis, acceleration by VE simulation, use of adjoints, ...
- Only a few simple examples presented herein
- Implemented as free open-source software: MRST-co2lab



Grid generation and coarsening

Discretization and solvers for incompressible flow

Discretization and solvers for compressible flow

Upscaling and multiscale methods

Fractured media and geomechanics

Workflow tools http://www.sintef.no/MRST